

The Global Weather Research and Forecasting (GWRf) Model: Model Evaluation, Sensitivity Study, and Future Year Simulation

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ABSTRACT

Global WRF (GWRf) is an extension of the mesoscale Weather Research and Forecasting (WRF) model that was developed for global weather research and forecasting applications. GWRf is being expanded to simulate atmospheric chemistry and its interactions with meteorology on a global scale. In this work, the ability of GWRf to reproduce major boundary layer meteorological variables that affect the fate and transport of air pollutants is assessed using observations from surface networks and satellites. The model evaluation shows an overall good performance in simulating global shortwave and longwave radiation, temperature, and specific humidity, despite large biases at high latitudes and over-Arctic and Antarctic areas. Larger biases exist in wind speed and precipitation predictions. These results are generally consistent with the performance of most current general circulation models where accuracies are often limited by a coarse grid resolution and inadequacies in sub-filter-scale parameterizations and errors in the specification of external forcings. The sensitivity simulations show that a coarse grid resolution leads to worse predictions of surface temperature and precipitation. The combinations of schemes that include the Dudhia shortwave radiation scheme or the Purdue Lin microphysics module, or the Grell-Devenyi cumulus parameterization lead to a worse performance for predictions of downward shortwave radiation flux, temperature, and specific humidity, as compared with those with respective alternative schemes. The physical option with the Purdue Lin microphysics module leads to a worse performance for precipitation predictions. The projected climate in 2050 indicates a warmer and drier climate, which may have important impacts on the fate and lifetime of air pollutants.

Keywords: Global Weather Simulation; Physical Options; Horizontal Grid Resolution

1. Introduction

The Weather Research and Forecasting (WRF) model has been developed by the National Center for Atmospheric Research (NCAR) to improve weaknesses of the Mesoscale Meteorological Model, Version 5 (MM5) and provide a flexible and portable open-source community model for both atmospheric research and operational forecasting [1,2]. The WRF system allows users to interchange various cores and physics packages, which is useful for inter-model evaluations and module sensitivity studies [3]. WRF has been utilized by thousands of users around the world in many different areas of atmospheric research including large-eddy simulations (e.g. [4]), real-time numerical weather predictions (NWP) (e.g. [5]), data assimilation [6,7], regional climate simulations (e.g. [3,8-12]), air quality modeling [e.g. 13-18], air quality

forecasting (e.g. [19-22]), and atmosphere-ocean coupling (e.g. [23,24]).

A global version of WRF (GWRf) released in 2008 is an extension of mesoscale WRF and a variant of planet WRF, which was initially designed to study the atmospheres and climate systems of other planets such as Mars, Titan, and Venus [25,26]. Four major modifications to the mesoscale WRF were made for application to the planetary atmosphere as the planet WRF: modification of the projection from an isotropic to a non-isotropic grid (*i.e.* to accommodate a latitude-longitude mesh), the addition of polar Fourier filters to remove model instabilities near the poles, adaptation of planetary constants and timing parameters, and parameterizations of sub-grid scale physical processes associated with specific planets [25]. To adapt the planetary model to the Earth, as opposed to other planets, certain Earth-specific planetary constants (e.g. acceleration due to gravity, reference

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pressure, and ideal gas constants) and necessary timing conventions (e.g. orbital parameters) have been incorporated into GWRP. GWRP enables modeling of global atmospheric circulation and the coupling between weather systems on global and regional scales with the same basic dynamics and physics [25,26]. Its initial evaluation showed an overall good performance in terms of the global zonal mean climatology for Earth, Mars, Titan, and Venus [26,27]. Compared to traditional general circulation models (GCMs) that have been developed since 1950s, GWRP enables a unified framework for the modeling of atmospheric processes and their interactions across scales spanning from global to local scales through 1-way or 2-way nesting. For example, GWRP can be used to provide initial conditions (ICs) and BCs for mesoscale WRF in nested simulations, which reduces inaccuracies arising from the use of different models with inconsistent model dynamics and physics.

While the mesoscale WRF model has been extensively evaluated using observations, there has been little evaluation of GWRP. In this work, GWRP version 3.0 is evaluated using available observations and reanalysis data. A number of sensitivity simulations are conducted to identify the most appropriate physical parameterizations in GWRP that produce the highest accuracy for global atmosphere. The main objectives of this study are to evaluate the capability of GWRP in reproducing global boundary layer meteorological variables that are most influential to air pollutants and to examine the sensitivity of the model predictions to various physical parameterizations. Such an evaluation is a critical step toward the extension of GWRP to include emissions and chemistry needed to simulate/forecast the global transport and fate of air pollutants, the impact of emissions on global air quality and radiative forcing, as well as forecasting the future climate change and its impact on air quality and *vice versa*.

2. Model Description and Simulation Design

GWRP provides a number of options for physical schemes or parameterizations [2]. The model radiation and physical schemes/parameterizations selected for the baseline and sensitivity simulations as well as a future year simulation are summarized in **Table 1**. The set of physics configurations in these simulations is selected with consideration of their suitability for long-term simulation over a global domain. For longwave radiation received at the surface (LW), the Community Atmosphere Model version 3 (CAM3.0) LW radiation scheme of Collins *et al.* [28] is used in the baseline simulation and the rapid radiative transfer model (RRTM) [29] is used in the sensitivity simulation. CAM3 LW is a spectral-band scheme with 2 bands used for climate simulations adapted from the NCAR CAM 3.0. It can handle water vapor, O₃, and CO₂, and interacts with model-resolved clouds and cloud fractions. RRTM LW is a spectral-band scheme with 16 bands using the correlated-*k* method, which calculates radiative transfer with *k* referring to the absorption coefficient. RRTM LW accounts for cloud optical depth and uses lookup tables to represent outgoing LW radiations caused by water vapor, O₃, CO₂, and trace gases. For shortwave radiation received at the surface (SW), the Goddard shortwave radiation scheme of Chou and Suarez [30] and Chou *et al.* [31] is used in the baseline simulation and the CAM3 shortwave radiation of Collins *et al.* [28] and the Dudhia scheme [32] are used in the sensitivity simulations. The Goddard SW scheme is a spectral band scheme with 11 bands. This scheme accounts for diffuse and direct solar radiation components in a two-stream approach, including the effects of water vapor, O₃, and CO₂ on radiation. The CAM3 SW is a spectral band that can handle several aerosol types and trace gases and their interactions with clouds. The CAM3 SW scheme is especially suited for

Table 1. Model configurations in the baseline and sensitivity simulations.

Atmospheric process	Parameterizations/schemes used in the baseline simulation	Parameterizations/schemes used in the sensitivity simulation
Longwave radiation	CAM [28]	The rapid radiative transfer model (RRTM) [29]
Shortwave radiation	Goddard scheme [30,31]	Dudhia [32] CAM [28]
Cloud microphysics	WSM3 [65,66]	Purdue Lin [33,34] WSM6 [65,66]
Land Surface Model	NOAH [36-38]	The simple thermal diffusion (Slab) scheme [39]
Cumulus parameterization	KF II [40-42]	GD ensemble [43]
PBL scheme	YSU scheme [44-46]	Same as baseline scheme
Surface-layer scheme	Monin-Obukhov scheme [47-51]	Same as baseline scheme

CAM—Community atmospheric model; GD—Grell-devenyi, RRTM—The rapid radiative transfer model; YSU—Yonsei University; WSM3—WRF single moment 3-class; NOAH—The National Center for Environmental Prediction (NNR), Oregon State University, Air Force, and Hydrologic Research Lab; and KF II—Kain-Fritsch version 2.

regional climate simulations [2]. The Dudhia SW scheme has a simple downward integration of solar flux and accounts for clear-air scattering, water vapor absorption, and cloud albedo and absorption.

For cloud microphysics parameterization, the WRF Single Moment 3-Class (WSM3) scheme is used in the baseline simulation and WSM 6-Class (WSM6) and Purdue Lin (PL) scheme [33,34] are used in the sensitivity simulations. WSM3 is a simple-ice scheme with three categories of hydrometers (*i.e.* water vapor, cloud water/ice, and rain/snow). It is computationally efficient, but does not treat super cooled water and gradual melting rates. WSM6 extends WSM3 to explicitly include water vapor, rain, snow, cloud ice, cloud water, and graupel. Among all WSM options in WRF, WSM6 is the most comprehensive option. It has been recently extended to a double moment warm rain microphysics (*i.e.* WDM6) [35]. The PL scheme treats six classes of hydrometeors: water vapor, cloud water, rain, cloud ice, snow, and graupel. For land-surface model (LSM), the National Center for Environmental Prediction (NCEP), Oregon State University, Air Force, and Hydrologic Research Lab (NOAH) model of Chen and Dudhia [36,37] and Ek *et al.* [38] is used in the baseline simulation and the simple thermal diffusion (Slab) scheme [39] is used in the sensitivity simulation. NOAH is a 4-layer soil temperature and moisture model with canopy moisture and snow cover prediction. The 4-layer thicknesses are 10, 30, 60 and 100 cm from the surface down. The NOAH scheme includes root zone, evapotranspiration, soil drainage, and runoff and accounts for vegetation categories, monthly vegetation fraction, soil texture, soil ice, fractional snow cover effects, surface emissivity properties, and improved urban treatment. The scheme provides sensible and latent heat fluxes to the boundary-layer scheme. It is used for both research and operational applications. The Slab LSM is based on the MM5 soil temperature model with 5 layers of 1, 2, 4, 8, and 16 cm thickness. The Slab LSM includes energy budget calculations accounting for radiation, sensible, latent heat flux, and a crude snow treatment with a constant snow cover. The soil moisture is fixed with a land-use and season-dependent constant value and vegetation effects are not explicitly considered. For the cumulus parameterization, the Kain-Fritsch II (KFII) [40-42] is used in the baseline simulation and the Grell-Devenyi (GD) ensemble [43] is used in the sensitivity simulation. The KF II scheme is a simple cloud model with moist updrafts and downdrafts and includes the effects of detrainment, entrainment, and simple microphysics. The GD scheme is an ensemble cumulus scheme which averages the results from multiple cumulus schemes with an equal weight. These cumulus schemes are all mass-flux type schemes, but with different updraft and downdraft entrainment and detrainment parameters,

and different precipitation efficiencies. In both baseline and sensitivity simulations, the Yonsei University (YSU) PBL scheme of Hong and Dudhia [44] and Hong *et al.* [45,46] and the Monin-Obukhov scheme adapted from MM5 [47] are used. The YSU scheme is a non-local scheme that includes an explicit treatment of the entrainment layer at the PBL top and provides a well-mixed boundary-layer profile. It also includes an enhanced stable boundary-layer diffusion algorithm that allows deeper mixing in windier conditions. The surface-layer scheme is based on the Monin-Obukhov similarity theory with Carlsion-Boland viscous sub-layer and standard similarity functions from look-up tables. It includes four stability regimes as described in Zhang and Anthes [48] and uses stability functions from Paulson [49], Dyer and Hicks [50], and Webb [51] to compute surface exchange coefficients for heat, moisture, and momentum.

Table 2 summarizes the configurations used in all simulations. GWRF is initialized using the WRF Preprocessing System version 3.0 (WPS3). The data used as input into WPS3 is from the NCEP Final Global Data Assimilation System (FNL), which has a horizontal grid resolution of 1° latitude \times 1° longitude and is available every six hours since July 1999. NCEP FNL is used to initialize the Global Forecasting System (GFS) model. To evaluate the performance of GWRF, simulations of the year 2001 are performed and analyzed. The simulations are initialized using the NCEP FNL analysis. The simulations are constrained by a monthly reinitialization of the WRF prognostic variables, including the soil temperature and skin temperature. Sea surface temperature (SST) and sea-ice fraction are reinitialized from weekly analyses, as recommended for simulations longer than one week [2]. A GWRF baseline simulation is conducted for the entire year of 2001 at a horizontal grid resolution of $1^\circ \times 1^\circ$, with 27 eta layers from 0 to 50 h Pa in the vertical domain. A number of sensitivity simulations are carried out to evaluate the impacts of physics parameterizations and grid resolutions on model performance. These sensitivity simulations include seven sets of simulations at $1^\circ \times 1^\circ$ (*i.e.* three for different combinations of longwave and shortwave radiation schemes (RAD1 for Dudhia SW and CAM3 LW, RAD2 for CAM3 SW and RRTM LW, and RAD3 for CAM3 SW and CAM3 LW), two for different cloud microphysics (CMP1 for WSM6 and CMP2 for PL), one for LSM (*i.e.* Slab), and one for cumulus parameterization (CCP) (*i.e.* GD)) and one with the same radiation and physics configurations as the baseline simulation but at a coarser horizontal grid resolution of 4° latitude \times 5° longitude.

To evaluate the capability of GWRF in simulating climate changes in future years, a simulation is completed for the year 2050 with the baseline physics configurations at a horizontal grid resolution of 4° latitude \times 5°

Table 2. Model configurations for GWRf baseline, sensitivity, and future year simulations.

Run Set	Run Name	Resolution	Time Period	Micro-Physics	Land-Surface Model	Cumulus Parameterization	Short-Wave	Long-Wave
Baseline	BASE	1° × 1°	2001	WSM3	NOAH	Kain-Fritsch	Goddard	CAM
	CMP1	1° × 1°	JJA	WSM6	NOAH	Kain-Fritsch	Goddard	CAM
	CMP2	1° × 1°	JJA	Purdue Lin	NOAH	Kain-Fritsch	Goddard	CAM
	LSM	1° × 1°	JJA	WSM3	SLAB	Kain-Fritsch	Goddard	CAM
	RAD1	1° × 1°	JJA	WSM3	NOAH	Kain-Fritsch	Dudhia	CAM
Sensitivity	RAD2	1° × 1°	JJA	WSM3	NOAH	Kain-Fritsch	CAM	RRTM
	RAD3	1° × 1°	JJA	WSM3	NOAH	Kain-Fritsch	CAM	CAM
	CCP	1° × 1°	JJA	WSM3	NOAH	Grell-Devenyi	Goddard	CAM
	OPT	1° × 1°	2001	WSM6	SLAB	Grell-Devenyi	CAM	RRTM
	Low_Res	1° × 1°	2001	WSM3	NOAH	Kain-Fritsch	Goddard	CAM
Future-Year	Future	1° × 1°	2050	WSM3	NOAH	Kain-Fritsch	Goddard	CAM

longitude. GWRf model outputs for 2050 are compared with results from the 2001 baseline simulation to analyze the variation trends of major meteorological variables. Since GWRf is not yet a climate model, it is initialized using atmospheric, land, and ocean/sea-ice output from the NCAR Community Climate System Model 3.0 (CCSM3) simulation for 2050 as part of the IPCC Special Report on Emission Scenario (SRES) B1 experiment. CCSM3 is a coupled atmosphere-ocean model at a spatial resolution of 257×129 grid points (1.4° latitude \times 1.4° longitude). The SRES B1 represents a low greenhouse gas concentration scenario with a CO₂ level of 550 ppm. Each month of the 2050 GWRf simulations were initialized using the corresponding monthly mean CCSM outputs. The 3-D variables used to initialize GWRf from the atmospheric component model of CCSM (*i.e.* Community Atmospheric Model) include vertical profiles of temperature (T), relative humidity (RH), geopotential height (Z3), the zonal (U) and meridional (V) components of wind speed, and surface pressure (PS). CAM outputs are vertically interpolated to map with the vertical structure of GWRf. Additional initialization data include 3-D soil temperature (TSOI) and moisture (H₂OSOI) predictions from the CCSM Community Land Model (CLM) and 2-D sea-ice fraction (ICEFRAC) from the CCSM Community Sea Ice Model (CSIM).

3. Evaluation Datasets and Methodology for 2001

3.1. Datasets for Model Evaluation

GWRf predictions are evaluated against surface observational networks and gridded reanalysis data which combine data from surface and satellite observations with other model outputs. A summary of the datasets is shown

in **Table 3**. The baseline surface radiation network (BSRN) is a global surface-based observational network established by the World Radiation Monitoring Center since 1992 for surface radiation fluxes at the Earth's surface for climate research [52]. It consists of 59 sites worldwide as of 2011, but only 28 sites have observations in 2001. The SW downward radiation flux measured by pyranometers and LW downward radiation flux measured by pyrgeometers every minute are used to compute hourly averaged observations for model evaluation. Meteorological parameters from the National Climatic Data Center (NCDC) are from the Global Climate Observing System (GSN) Surface Network, Monthly (GSNMON), with over 900 sites worldwide. The observed mean monthly temperature (°C) and total monthly precipitation (mm) for ~200 sites in 2001 are used for model evaluation. The Global Precipitation Climatology Project (GPCP) is a part of the Global Energy and Water Cycle Experiment (GEWEX) of the World Climate Research program (WCRP). The monthly mean precipitation data are produced from an analysis designed by the Global Precipitation Climatology Centre by merging precipitation estimates from microwave, infrared, and sounder data from international precipitation-related satellites, as well as from precipitation gauges over 6000 stations on a $2.5^\circ \times 2.5^\circ$ latitude-longitude grid from 1979 to present [53]. The version of GPCP used in this study is the 2.5-degree version 2. The NCEP and the National Center for Atmospheric Research (NCAR) Reanalysis (NNR) dataset is widely used to validate global simulations in the atmospheric modeling community. Kalnay *et al.* [54] published a 40-year record (1957-1996) of global atmospheric parameters from reanalysis. The reanalysis uses a combination of assimilated observations (remotely-sensed and surface based) at a horizontal grid

Table 3. Summary of datasets used to evaluate GWRP.

Database ^a	Variables ^b	Data frequency	Number of sites as of 2001
BSRN	SW, LW	Minutely	~28
NCDC	T2, Precip	Monthly	~200
GPCP	Precipitation	Monthly	Global domain, Gridded (2.5° × 2.5°)
NNR Reanalysis (NNR)	T2, Q2, U10, V10, SW, LW	4-times daily	Global domain, Gridded (2.5° × 2.5°)

^aBSRN—Baseline Surface Radiation network, <http://www.gewex.org/bsrn.html>; NCDC—National Climatic Data Center, <http://www.ncdc.noaa.gov/oa/ncdc.html>; GPCP—Global Precipitation Climatology Project, <http://precip.gsfc.nasa.gov> and Xie and Arkin, 1997; NNR Reanalysis (NNR), <http://www.NNR.noaa.gov>; ^bSW—Downward shortwave flux; LW—Downward longwave flux; T2—Temperature at 2-meter; Precip—Precipitation; Q2—Water vapor mixing ratio at 2-meter; U10—Zonal mean wind speed at 10-m; V10—Meridional mean wind speed at 10-m.

resolution of 2.5° × 2.5° latitude/longitude for each year and model regression to produce global analyses of atmospheric fields.

3.2. Evaluation Protocol

The model evaluation focuses on major boundary layer meteorological variables including a combination of non-convective and convective weekly accumulated precipitation (RAIN + RAINNC), 2-meter temperature (T2) and specific humidity (Q2), and 10-meter wind velocities and their zonal (U10) and meridional (V10) components, as well as radiation variables such as SW and LW radiation. The overall performance of GWRP is evaluated in terms of spatial distribution, seasonal and temporal variations, and statistics over the global domain, the Northern and Southern Hemispheres, and the six populated continents. The six circulations cells are also used as subdomains for model evaluation, they include: the Polar (60°N - 90°N and 60°S - 90°S), Ferrel (30°N - 60°N and 30°S - 60°S), and Hadley (0°N - 30°N and 0°S - 30°S) Cells in the Northern and Southern hemispheres. The statistical measures include normalized mean bias (NMB), normalized mean error (NME), mean bias (MB), root mean square error (RMSE), and correlation coefficient (Corr) over the entire domain, sub-domains, and continental domains. The formulas used to calculate these statistical metrics are taken from Zhang *et al.* [55].

4. Evaluation of Baseline Results

Table 4 summarizes the overall performance statistics for all meteorological variables. **Figures 1** and **2** show simulated downward SW and LW radiation fluxes overlaid with observations from BSRN during winter and summer. GWRP reproduces the radiation observations reasonably well in both seasons with moderate overpredictions (MBs of 38.5 - 51.5 W·m⁻² and NMBs of 24% - 27%) for SW radiation and underproductions (MBs of -32.1 to -30.0 W·m⁻² and NMBs of -10.3% to -9.2%) for LW radiation. As shown in **Table 4** and **Figure 3**, similar overpredictions of SW and underpredictions of

LW radiation fluxes against the NNR data also occur on a global scale but to a lesser extent than against the BSRN data. Such overpredictions of SW fluxes and underpredictions of LW radiation occur mainly between 30°N and 30°S and dominate the trends of the annual predictions with MBs of 16.6 W·m⁻² for SW and -16.6 W·m⁻² for LW radiation against the NNR data. These results are overall consistent with other GCMs. Wild *et al.* [56] reported that a tendency common to all GCMs which overestimate SW by an average of 10 - 15 W·m⁻², due likely to an underestimation of atmospheric absorption. The performance of ECHAM3 in terms of SW radiation also showed a latitudinal dependence, with overpredictions by up to 40 W·m⁻² over the low latitudes and underestimations by up to 20 W·m⁻² over the high latitudes due to an underestimation of cloud cover in the annual mean. Annual performance statistics for GWRP compares well with these trends, exhibiting the largest overpredictions for SW radiation fluxes over the Northern (MB of 41.1 W·m⁻² and NMB of 16.8%) and Southern (42.9 W·m⁻² and NMB of 17.9%) Hadley Cells, and the largest underpredictions over the North Pole and North Ferrel cells which have MBs of -9.0 W·m⁻² and -3.3 W·m⁻² and NMBs of -6.8% and -1.7 W·m⁻², respectively (see **Figure 3** for NMBs). Consistently, GWRP also shows large overpredictions of SW in major populated continents over the low latitudes such as South America, Australia, and Africa and underpredictions for the continent over the high latitude such as Europe in terms of winter, summer, and annual mean values (see **Figure 3**). Wild *et al.* [56] indicated a tendency of GCMs to underestimate LW by an average of 10 - 15 W·m⁻² (corresponding to the global mean observed and simulated values of 293 W·m⁻² and 278 W·m⁻², respectively), due to an underestimation of low-level clouds. For comparison, GWRP underestimates LW radiation with a global annual MB value of -16.6 W·m⁻² (corresponding to mean observed and simulated values of 295 W·m⁻² and 278 W·m⁻², respectively). It also shows large underpredictions of LW fluxes in all six major populated

Table 4. Performance statistics of GWRf at a horizontal grid resolution of 1° × 1° on a global scale.

Variable	Dataset	Season	Mean Obs	Mean Mod	Number	Corr	MB	RMSE	NMB (%)	NME (%)
SW ($W \cdot m^{-2}$)	NNR	Annual	186.41	203.02	64,080	0.93	16.62	32.40	8.91	12.91
	NNR	Summer	191.34	198.94	61,560	0.93	7.59	41.34	3.97	15.08
	BSRN	Summer	218.83	270.32	29	0.98	51.50	57.07	23.53	23.53
	NNR	Winter	205.55	229.08	61547	0.96	23.53	41.55	11.45	14.04
	BSRN	Winter	143.08	181.61	28	0.98	38.53	46.42	26.92	27.08
LW ($W \cdot m^{-2}$)	NNR	Annual	295.03	278.43	64,080	0.99	-16.60	23.54	-5.63	6.75
	NNR	Summer	306.45	293.76	64,080	0.98	-12.68	24.30	-4.14	6.65
	BSRN	Summer	350.01	317.91	29	0.93	-32.10	39.63	-9.17	9.17
	NNR	Winter	287.84	267.24	64,080	0.98	-20.59	26.92	-7.15	8.09
	BSRN	Winter	291.18	261.34	28	0.93	-30.04	39.59	-10.32	10.79
T2 ($^{\circ}C$)	NNR	Annual	5.38	5.36	64,080	1.00	-0.02	1.83	-0.38	21.34
	NNR	Summer	7.48	7.68	64,079	0.99	0.20	2.33	2.70	18.75
	NCDC	Summer	20.77	20.27	199	0.88	-0.51	4.18	-2.43	12.24
	NNR	Winter	4.15	3.55	64,080	0.99	-0.60	2.55	-14.52	37.52
	NCDC	Winter	11.62	9.62	243	0.55	-2.00	24.75	-17.21	36.23
Q2 ($g \cdot kg^{-1}$)	NNR	Annual	8.17	7.98	64,080	0.99	-0.19	0.86	-2.29	6.39
	NNR	Summer	8.85	8.64	64,080	0.99	-0.21	1.03	-2.39	7.12
	NNR	Winter	7.72	7.55	64,080	0.99	-0.17	0.83	-2.22	6.42
U10 ($m \cdot s^{-1}$)	NNR	Annual	-0.01	-0.06	64,080	0.96	-0.05	0.97	-452.13	6471.81
	NNR	Summer	0.04	-0.12	64,080	0.93	-0.16	1.41	-387.21	2554.53
	NNR	Winter	-0.06	-0.10	64,080	0.94	-0.04	1.34	-68.47	1642.83
V10 ($m \cdot s^{-1}$)	NNR	Annual	0.15	0.11	64,080	0.90	-0.04	0.85	-26.16	423.01
	NNR	Summer	0.49	0.47	64080	0.87	-0.02	1.24	-3.95	189.73
	NNR	Winter	-0.17	-0.22	64,078	0.86	-0.05	1.13	-31.31	509.57
Precip ($mm \cdot d^{-1}$)	GPCP	Annual	2.15	2.82	10,368	0.9	0.67	1.47	31.03	40.06
	GPCP	Summer	2.22	3.08	10,363	0.83	0.86	1.96	38.64	54.4
	NCDC	Summer	93.20	116.77	172	0.67	23.57	88.50	25.29	58.48
	GPCP	Winter	2.09	2.39	10355	0.83	0.3	1.7	14.39	44.33
	NCDC	Winter	82.83	78.56	223	0.71	-4.27	80.97	-5.15	55.36

SW—Downward shortwave radiative flux at surface; LW—Downward longwave flux at surface; T2—Temperature at 2-m; Q2—Water vapor mixing ratio at 2-m; U10—Zonal mean wind speed at 10-m; V10—Meridional mean wind speed at 10-m; Precip—Daily mean precipitation rate.

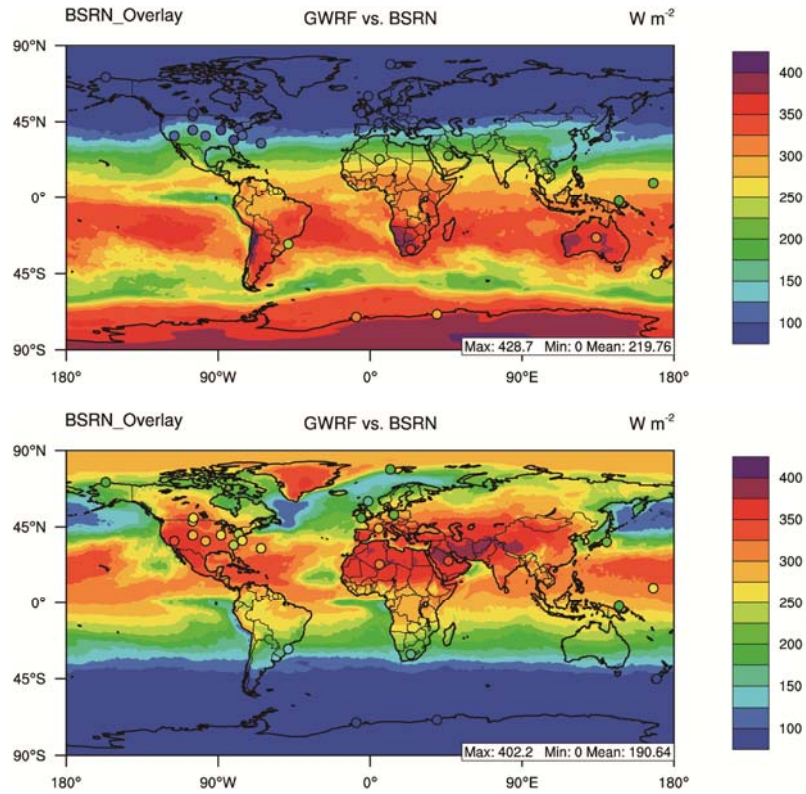


Figure 1. 2001 winter (top) and summer (bottom) downward shortwave fluxes at the surface simulated at the horizontal grid resolution of 1° × 1° overlaid with observations from BSRN (denoted by circles).

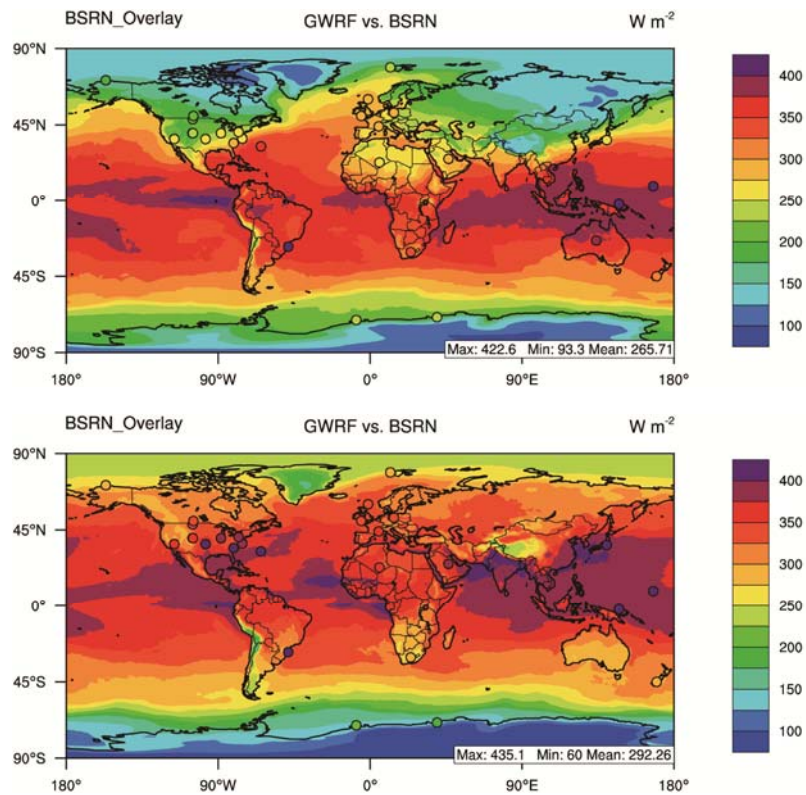


Figure 2. 2001 winter (top) and summer (bottom) downward longwave fluxes at the surface simulated at the horizontal grid resolution of 1° × 1° overlaid with observations from BSRN (denoted by circles).

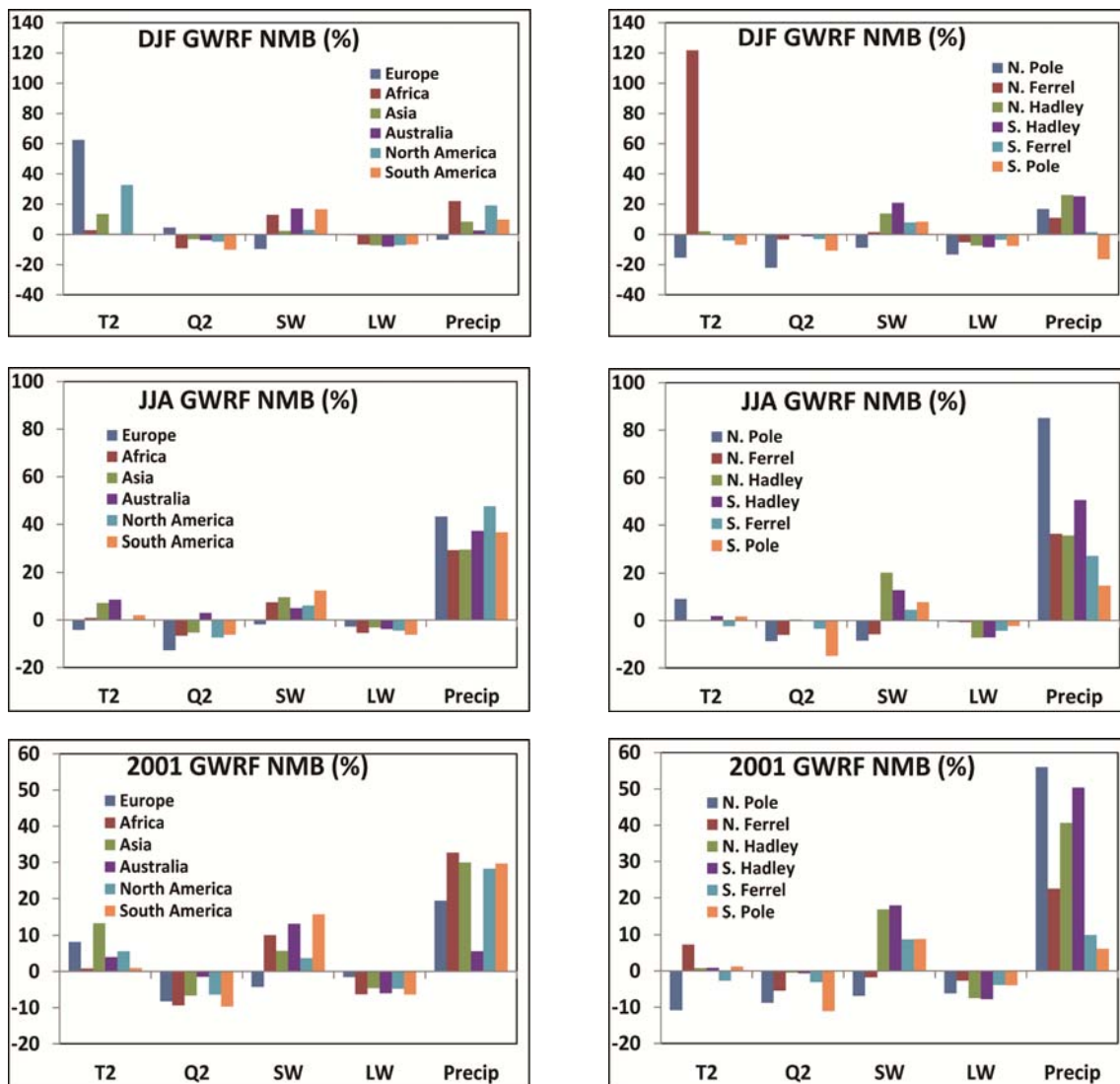


Figure 3. 2001 winter (top), summer (middle), and annual (bottom) mean normalized mean bias (NMB) (%) of T2, Q2, SW, and LW against the NNR data and Precip against the GPCP data at a horizontal grid resolution of $1^\circ \times 1^\circ$ for the six continental domains (left column) and the six circulation cell domains (right column).

continents and circulation cells in seasonal and annual mean values (see **Figure 3**).

Figure 4 shows simulated T2 overlaid with observations from NCDC during winter and summer. GWRP generally reproduces well the spatial distribution of observed T2 in both seasons with mean cold bias of -2.0°C in winter and -0.5°C in summer (NMBs of -17.2% and -2.4% , respectively). The largest cold biases occur in the North Pole in winter and in the South Ferrell cell in summer. A similar trend for T2 is found against the NNR data (see **Figure 3**). Over the North Ferrell cell, the mean NNR and simulated values are -0.59°C and 0.13°C during summer, respectively, resulting in an MB of 0.72°C and NMB of 121.8% , as shown in **Figure 3**. Such an overprediction is caused by overestimations in SW radiation fluxes over the Northern mid-latitudes. For major con-

tinents, an overprediction of SW radiation fluxes correlates well with an overprediction of T2 (except for Europe). The anti-correlation exists between T2 and SW over Europe in the summer, winter, and annual means and over the North Ferrel cell in the annual means. Although SW is underpredicted over Europe during winter, the simulated T2 of -1.0°C is warmer than the NNR value of -2.7°C , leading to a high NMB of 62.6% in winter and an annual NMB of 8.2% . Similarly, during winter the simulated T2 of 0.13°C is much warmer than the observed T2 of -0.59°C , leading to a very high winter NMB of 121.8% and an annual NMB of 7.3% over the Northern Ferrel cell. Large errors associated with surface temperature exist over the high latitudes and areas of sharp terrain gradients, which are consistent with most GCMs.

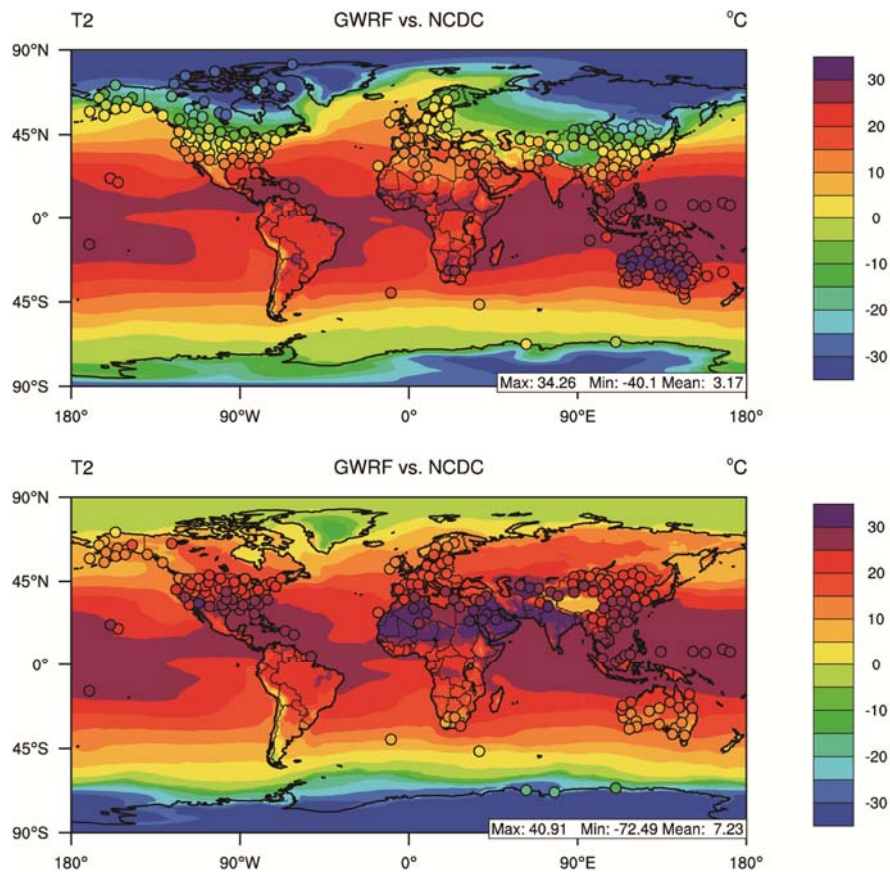


Figure 4. 2001 winter (top) and summer (bottom) temperatures at 2-meter (T2) simulated at the horizontal grid resolution of $1^\circ \times 1^\circ$ overlaid with observations from NCDC (denoted by circles).

For example, Christensen *et al.* [57] reported similar areas of the globe where models are known to perform poorly for surface temperature such as the Tibetan Plateau, Arctic, and Antarctic. According to the 4th Assessment Report (AR4) of the Intergovernmental Panel on Climate Change (IPCC), over most of the globe, the simulated annual mean surface temperature by most current GCMs differs from observations by less than 2°C [58,59]. Individual models typically have larger errors (but in most cases still less than 3°C , except at high latitudes) in regions of sharp elevation changes as a result of inconsistencies between simulated and actual topography [59,60].

The expected performance for GCMs for specific humidity is to have a bias of $<1 \text{ g}\cdot\text{kg}^{-1}$ suggested by Lambert and Boer, (2001) or error of $<10\%$ suggested by Randall *et al.* [60]. As shown in **Table 4**, GWRP gives a good agreement of Q2 against the NNR data, with domain wide annual MB of $-0.2 \text{ g}\cdot\text{kg}^{-1}$ and NMB of -2.3% . The largest underpredictions occur in the North and South Poles among the six circulation cells and South America and Africa among the six populated continents. The largest underestimations occur over humid regions (e.g. areas of India and sub-Saharan Africa), and the

largest overestimations occur over shallow areas of water that are exposed to a large amount of insolation year round (e.g. Gulf of California, Persian Gulf, and Red Sea). As shown in **Figure 3**, an underprediction of Q2 correlates with an underprediction of LW radiation over all regions annually, which is physically consistent because LW is a function of the vertical distribution of atmospheric absorbers, of which a major constituent is water vapor. As shown in **Table 4**, GWRP gives a relatively good agreement for V10 but a much worse one for U10 against the NNR data, with domain wide annual MB of $-0.04 \text{ m}\cdot\text{s}^{-1}$ and NMB of -26.2% for V10 and MB of $-0.05 \text{ m}\cdot\text{s}^{-1}$ and NMB of -452.1% for U10. The very high NMB for U10 is caused by a very low U10 value of $-0.01 \text{ m}\cdot\text{s}^{-1}$ from the NNR data.

Figures 5 and 6 show simulated precipitation overlaid with monthly-mean precipitation from NCDC and the mean bias against the daily-mean precipitation GPCP, respectively, during winter and summer. Compared to the NCDC observations, GWRP moderately overpredicts precipitation in summer and slightly underpredicts it in winter, with NMBs of 25.3% and -5.2% , respectively. GWRP generally reproduces the spatial distributions of observed precipitation in both seasons and some regional

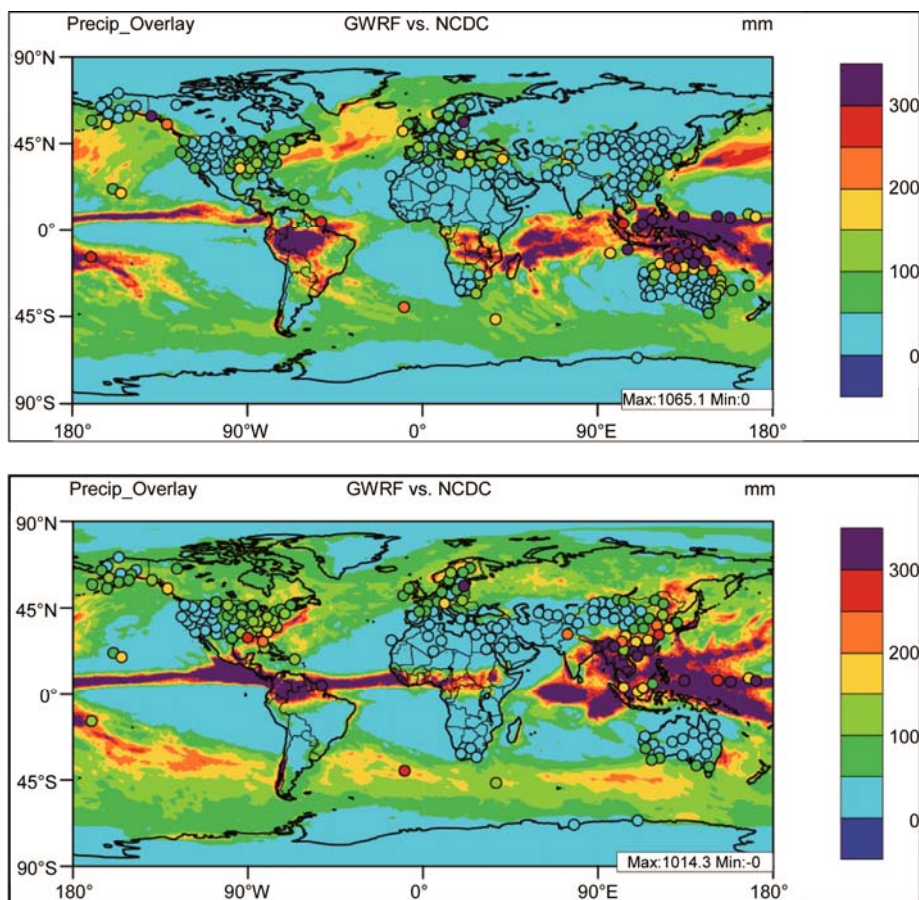


Figure 5. 2001 winter (top) and summer (bottom) monthly mean accumulated precipitation simulated at a horizontal grid resolution of $1^\circ \times 1^\circ$ overlaid with observations from NCDC (denoted by circles).

variations (e.g. the high precipitation areas in the southeastern U.S.) during summer. However, it significantly underpredicts at a few sites (e.g. in southern U.S., Middle East, Northern Australia, and over the oceans during winter and in western Canada, northern India, and over the oceans during summer). Compared to the GPCP daily precipitation data of Xie and Arkin [53], moderate overpredictions occur over the global domain in terms of winter, summer, and annual mean values, with MBs of 0.3 to $0.9 \text{ mm} \cdot \text{day}^{-1}$ and NMBs of 14.4% to 38.6%. As shown in **Figure 6**, the largest overpredictions occur over mid-latitudes during winter and mid-latitudes and the South Ferrel Cell during summer. Randall *et al.* [60] indicated that the rain-bearing systems used in establishing mean precipitation climatology are the largest source of inaccuracy for precipitation in GCMs, which predict too frequent rain but with reduced intensity. Compared to GPCP data of Xie and Arkin [53] and the observations of Randall *et al.* [60], GWRP overpredicts rainfall over mid-latitudes of both the Northern and Southern Hemispheres and underestimates precipitation in the tropical convergence zones. Similar to most GCMs, GWRP performs poorly for both temperature and precipitation in

several areas including the Tibetan Plateau, Arctic, and Antarctic, due possibly to difficulties simulating the effects of topography and albedo feedbacks under a heavy snow cover, biased atmospheric storm tracks and sea ice cover, the use of a coarse resolution for topography and simplified parameterizations, and the differences among the models in their sub-grid scale parameterization schemes [57,61].

5. Sensitivity Study

5.1. Sensitivity to Horizontal Grid Resolution

Figure 6 shows that the precipitation predictions at a horizontal grid resolution of $4^\circ \times 5^\circ$ are much worse than those at a finer scale, particularly in the Tropics (Northern and Southern Hadley Cells) in both seasons and annually. The performance statistics against GPCP precipitation data in **Tables 4** and **5** shows much larger overpredictions with NMBs of 42.8%, 52.7%, and 49.4% at $4^\circ \times 5^\circ$ versus NMBs of 14.4%, 38.6%, and 31.0% at $1^\circ \times 1^\circ$ for winter, summer, and annual mean values, respectively. **Figure 7** shows NMBs of major meteorological variables including T2, Q2, SW, and LW against the NNR data

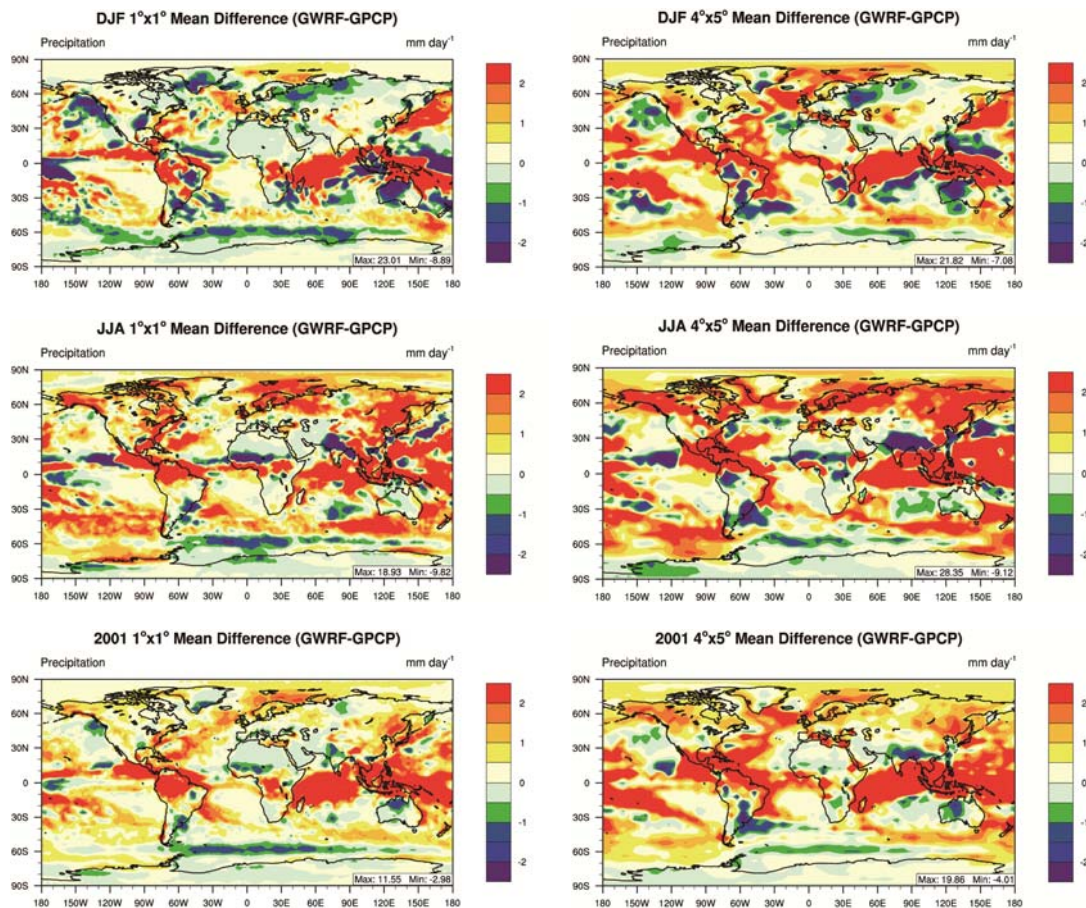


Figure 6. 2001 winter (top), summer (middle), and annual (bottom) mean bias plot of daily mean precipitation rate against GPCP data at a horizontal grid resolution of $1^\circ \times 1^\circ$ (left column) and $4^\circ \times 5^\circ$ (right column).

and Precip against the NCDC data at a horizontal grid resolution of $4^\circ \times 5^\circ$ over major continents and circulation cells. The corresponding performance statistics for these variables and additional variables such as U10 and V10 are summarized in **Table 5**. The spatial distributions of summer mean biases for SW, T2, and Precip at both $1^\circ \times 1^\circ$ and $4^\circ \times 5^\circ$ (simulations BASE and Low_Res, respectively) are also compared in **Figures 8-10**. Compared with those at $1^\circ \times 1^\circ$ (see **Figure 3** and **Table 4**), the model predictions of SW and LW radiation fluxes are similar, with a slightly better performance at $4^\circ \times 5^\circ$. As shown in **Figure 8**, the model at $4^\circ \times 5^\circ$ shows smaller overpredictions of SW radiation flux over tropics but larger underpredictions at higher latitudes in the Northern Hemisphere. As shown in **Figure 9**, the model at $4^\circ \times 5^\circ$ shows smaller underpredictions of LW radiation flux over tropics. As shown in **Figures 3, 7, and 10**, the model performs similarly for T2 in summer and on an annual basis at both grid resolutions, except that T2 values over North America and North Pole change from overpredictions to underpredictions in summer. However, the model shows a larger sensitivity during winter, with a much worse overprediction over North America

and a much less overprediction over North Ferrel cell. Q2 is relatively insensitive to horizontal grid resolution. The model at $4^\circ \times 5^\circ$ gives a larger underprediction for U10 in the winter and annual mean but better predictions for V10 in winter and annual mean. A large sensitivity to the horizontal grid resolution is found for Precip. A coarser resolution causes worse agreement with the NCDC data for Precip over all major continents and most circulation cells during winter and most continents (e.g., Europe, North America, Africa, Asia) and circulation cells (e.g. North Pole, North Ferrel, North Hadley, and South Hadley) during summer, as shown in **Figures 7 and 11**. On an annual basis, a coarser resolution causes worse agreement with the NCDC's data for Precip over most continents (e.g. Europe, North America, Africa, Asia) and all circulation cells. These results are consistent with performance statistics of Precip against the NCDC data in **Table 5** and its spatial distribution of mean biases against the GPCP data in **Figure 6**.

5.2. Sensitivity to Physical Parameterizations

In addition to model results at different grid resolutions,

Table 5. Performance statistics of GWRF at a horizontal grid resolution of $4^\circ \times 5^\circ$ on a Global scale.

Variable	Dataset	Season	Mean Obs	Mean Mod	Number	Corr	MB	RMSE	NMB (%)	NME (%)
SW ($\text{W}\cdot\text{m}^{-2}$)	NNR	Annual	185.88	194.29	3240	0.93	8.41	25.40	4.52	10.75
	NNR	Summer	191.94	191.55	3096	0.94	-0.39	38.20	-0.20	13.53
	BSRN	Summer	218.83	264.83	29	0.97	46.00	54.10	21.03	21.02
	NNR	Winter	206.68	220.93	3096	0.96	14.26	35.87	6.90	12.04
	BSRN	Winter	143.08	173.31	28	0.98	30.23	39.25	21.13	22.75
LW ($\text{W}\cdot\text{m}^{-2}$)	NNR	Annual	293.27	280.66	3240	0.99	-12.61	21.59	-4.30	6.12
	NNR	Summer	304.77	293.98	3240	0.98	-10.79	23.48	-3.54	6.52
	BSRN	Summer	350.01	316.06	29	0.89	-33.95	44.34	-9.70	10.50
	NNR	Winter	286.15	271.41	3240	0.98	-14.74	24.93	-5.15	7.35
	BSRN	Winter	291.18	267.82	28	0.91	-23.37	36.63	-8.02	9.51
T2 ($^\circ\text{C}$)	NNR	Annual	4.96	4.96	3240	0.99	0.00	2.33	0.07	26.76
	NNR	Summer	7.07	6.87	3240	0.99	-0.19	2.90	-2.73	23.84
	NCDC	Summer	20.77	19.12	199	0.84	-1.66	5.12	-7.99	17.02
	NNR	Winter	3.80	3.46	3239	0.99	-0.34	2.84	-8.87	46.51
	NCDC	Winter	11.62	9.32	243	0.54	-2.30	80.97	-19.77	42.29
Q2 ($\text{g}\cdot\text{kg}^{-1}$)	NNR	Annual	8.09	7.87	3240	0.99	-0.22	1.04	-2.67	6.83
	NNR	Summer	8.78	8.45	3240	0.98	-0.32	1.29	-3.67	8.03
	NNR	Winter	7.65	7.49	3240	0.99	-0.16	1.08	-2.08	7.31
U10 ($\text{m}\cdot\text{s}^{-1}$)	NNR	Annual	-0.02	-0.17	3240	0.93	-0.15	1.37	-826.26	5851.79
	NNR	Summer	0.04	-0.10	3240	0.91	-0.14	1.68	-377.97	3460.50
	NNR	Winter	-0.07	-0.30	3240	0.88	-0.23	1.82	-332.73	1985.28
V10 ($\text{m}\cdot\text{s}^{-1}$)	NNR	Annual	0.15	0.13	3240	0.85	-0.02	1.06	-13.28	507.30
	NNR	Summer	0.49	0.45	3240	0.84	-0.04	1.38	-7.94	202.45
	NNR	Winter	-0.16	-0.13	3239	0.78	0.03	1.35	18.66	641.77
Precip ($\text{mm}\cdot\text{d}^{-1}$)	GPCP	Annual	2.15	3.22	3240	0.82	1.06	2.29	49.37	59.32
	GPCP	Summer	2.22	3.39	3240	0.74	1.17	2.72	52.66	72.09
	NCDC	Summer	93.20	125.88	172	0.64	32.68	101.87	35.07	70.98
	GPCP	Winter	2.08	2.98	3240	0.78	0.89	2.45	42.83	63.59
	NCDC	Winter	82.83	88.99	223	0.74	6.16	79.05	7.45	57.31

SW—Downward shortwave radiative flux at surface; LW—Downward longwave flux at surface; T2—Temperature at 2-m; Q2—Water vapor mixing ratio at 2-m; U10—Zonal mean wind speed at 10-m; V10—Meridional mean wind speed at 10-m; Precip—Daily mean precipitation rate.

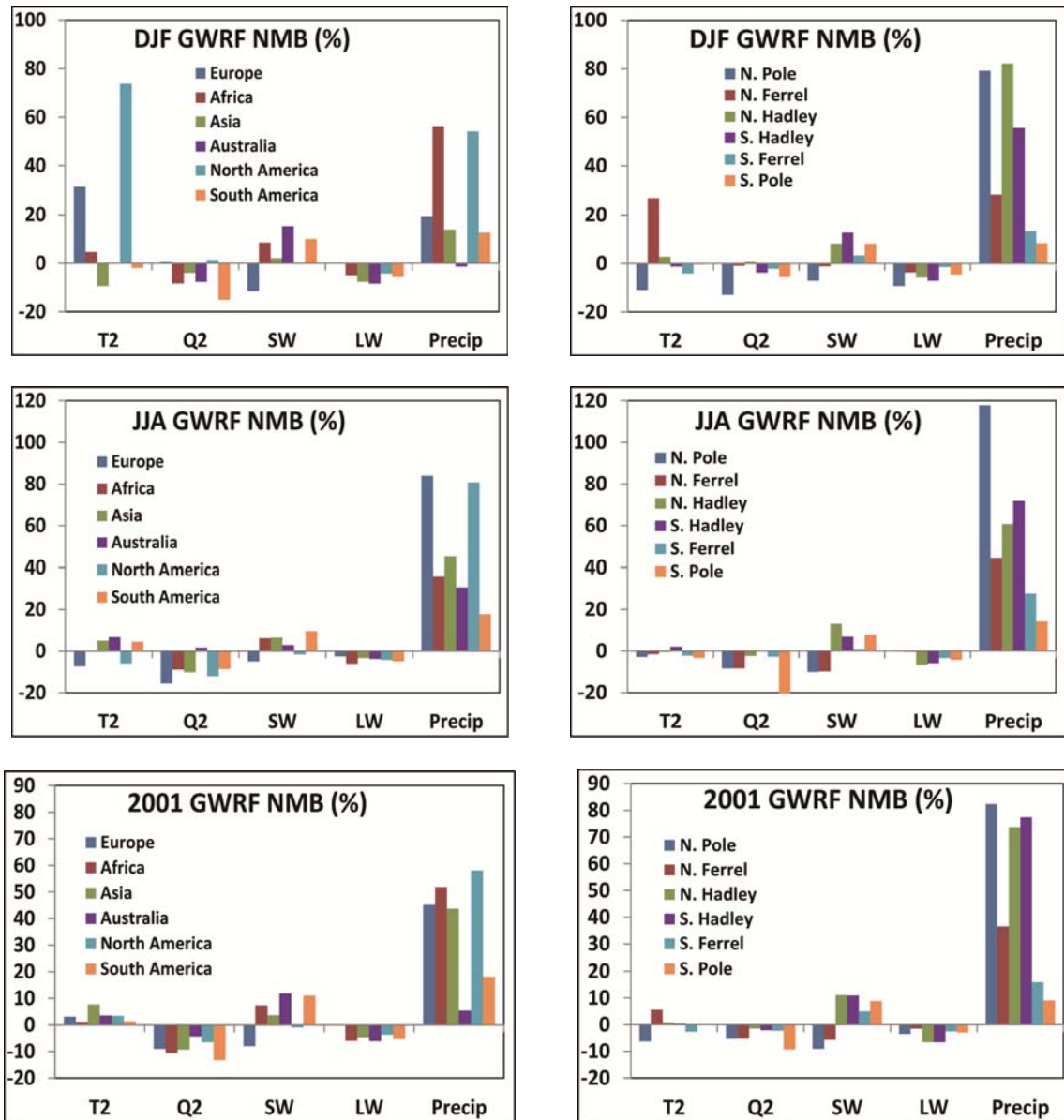


Figure 7. 2001 winter, summer, and annual mean normalized mean bias (NMB) (%) of T2, Q2, SW, and LW against the NNR data and Precip against the GPCP data at a horizontal grid resolution of $4^\circ \times 5^\circ$ for the six continental domains (left column) and the six circulation cell domains (right column).

Figures 8-11 also compare the spatial distributions of summer mean biases of SW, T2, and Precip at the surface against the NNR data at a horizontal grid resolution of $1^\circ \times 1^\circ$ from baseline simulations and sensitivity simulations including RAD1, RAD2, RAD4, CMP1, CMP2, LSM, and CCP. Table 6 summarizes model performance statistics against available observations for SW, LW, T2, Q2, and Precip with various physical options for summer 2001.

As shown in Figure 8, SW is overestimated in the tropics and subtropics ($30^\circ\text{N} - 30^\circ\text{S}$) by all simulations except for CCP, with relatively smaller biases in the simulations CMP2 and RAD1. Large underpredictions

occur at high latitudes in the Northern Hemisphere for all simulations (except for RAD2), particularly in RAD1, indicating a model deficiency in radiation involving ice. Underpredictions also occur over areas off the west coasts of Africa and South America in the South Hadley Cell in all runs except for RAD2 and in the South Ferrel Cell for RAD1 and CCP. In addition, CCP gives a large underestimation over all ocean surfaces outside of the Polar Regions. Among all sensitivity simulations, LSM gives the best performance in terms of statics, with an MB of $2.1 \text{ W}\cdot\text{m}^{-2}$ and an NMB value of 1.1%, and RAD1 gives the worst performance, with an MB of $-31.7 \text{ W}\cdot\text{m}^{-2}$ and an NMB value of -16.6%. The NMB values are

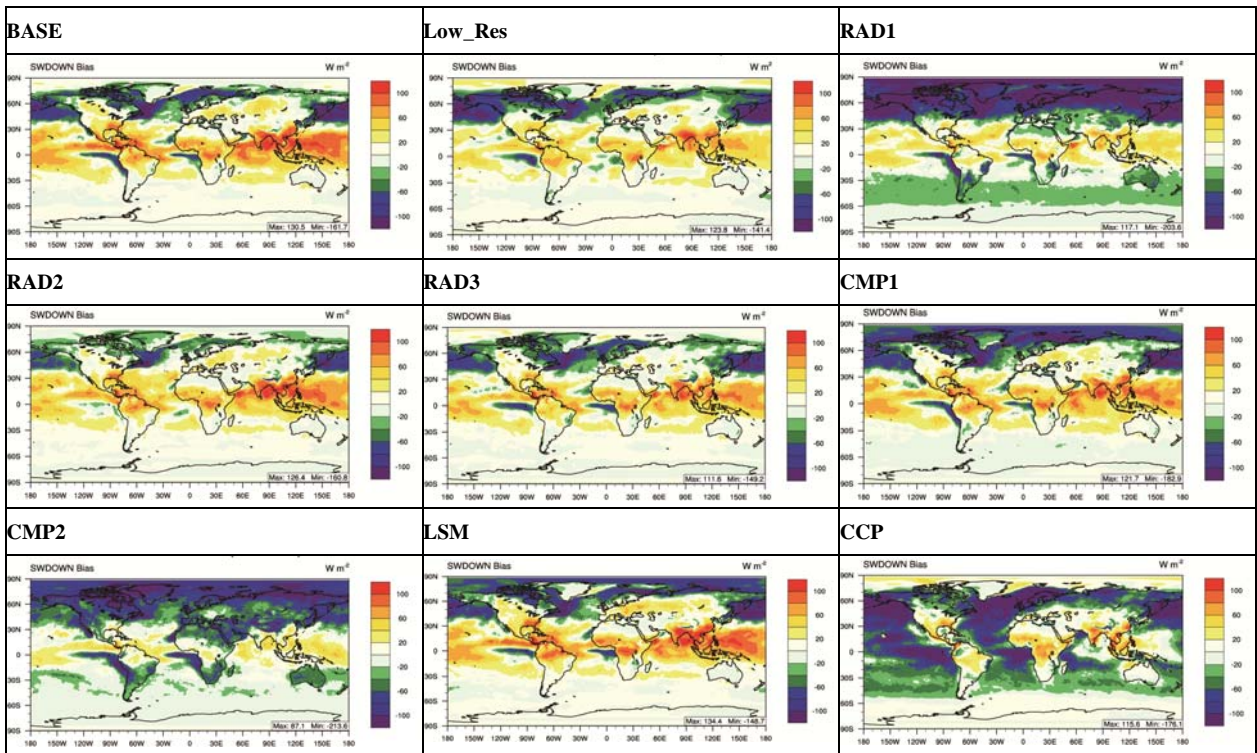


Figure 8. 2001 summer mean bias plot of downward shortwave fluxes at the surface against NNR data at a horizontal grid resolution of $1^\circ \times 1^\circ$ from the baseline, RAD1, RAD2, RAD4, CMP1, CMP2, LSM, and CCP simulations and at a horizontal grid resolution of $4^\circ \times 5^\circ$ from the Low_Res simulation.

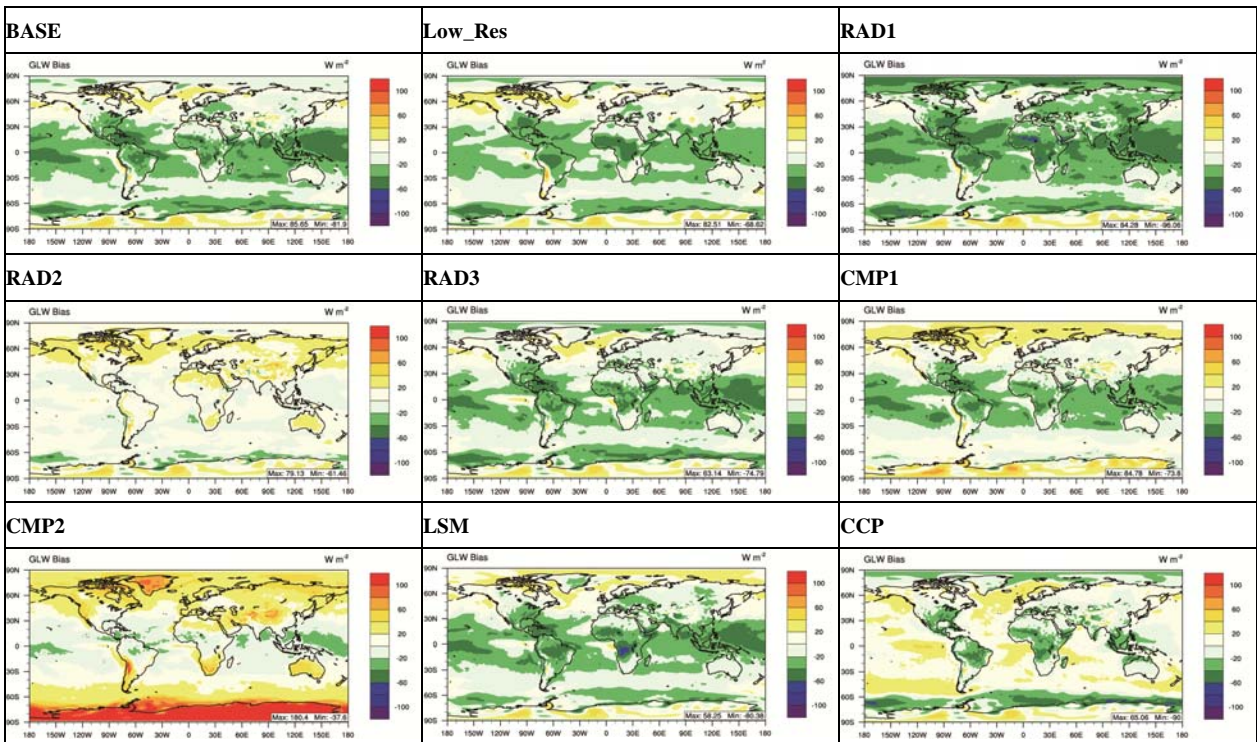


Figure 9. 2001 summer mean bias plot of downward longwave fluxes at the surface against NNR data at a horizontal grid resolution of $1^\circ \times 1^\circ$ from the baseline, RAD1, RAD2, RAD4, CMP1, CMP2, LSM, and CCP simulations and at a horizontal grid resolution of $4^\circ \times 5^\circ$ from the Low_Res simulation.

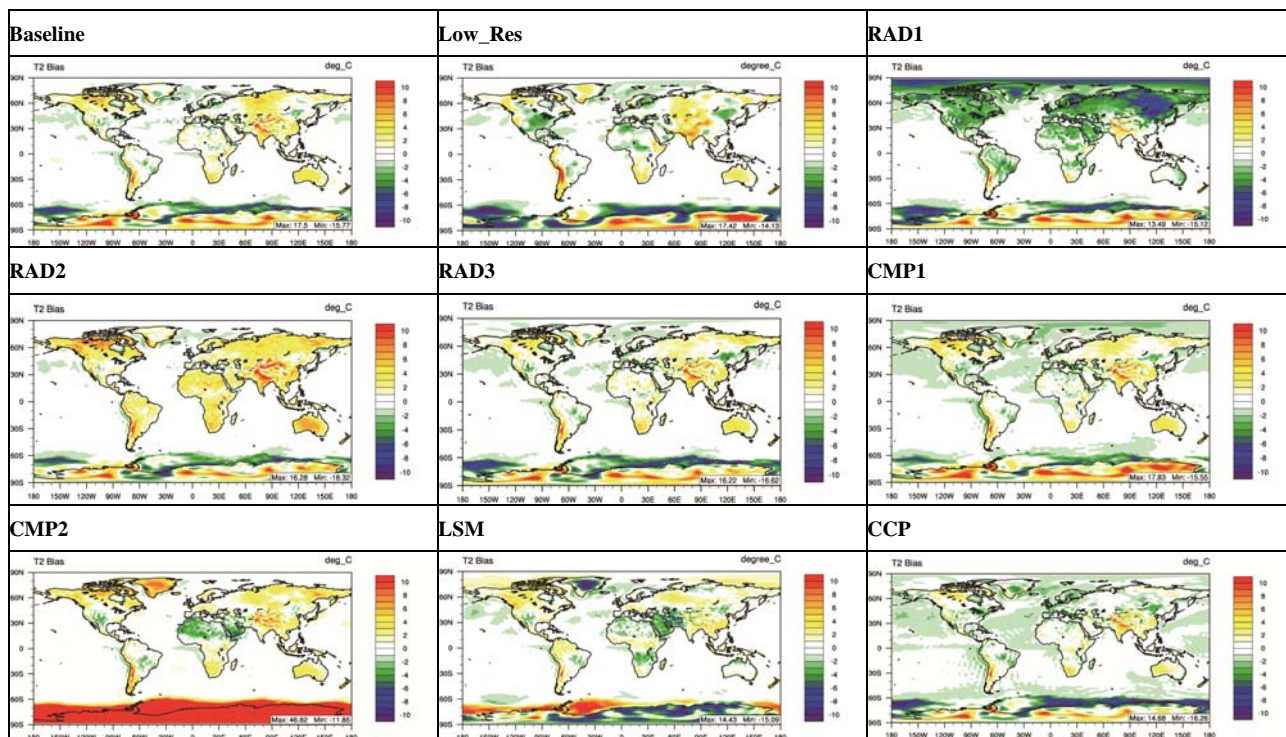


Figure 10. 2001 summer mean bias plot of temperatures at 2-m against NNR data at a horizontal grid resolution of $1^\circ \times 1^\circ$ from the baseline, RAD1, RAD2, RAD4, CMP1, CMP2, LSM, and CCP simulations and at a horizontal grid resolution of $4^\circ \times 5^\circ$ from the Low_Res simulation.

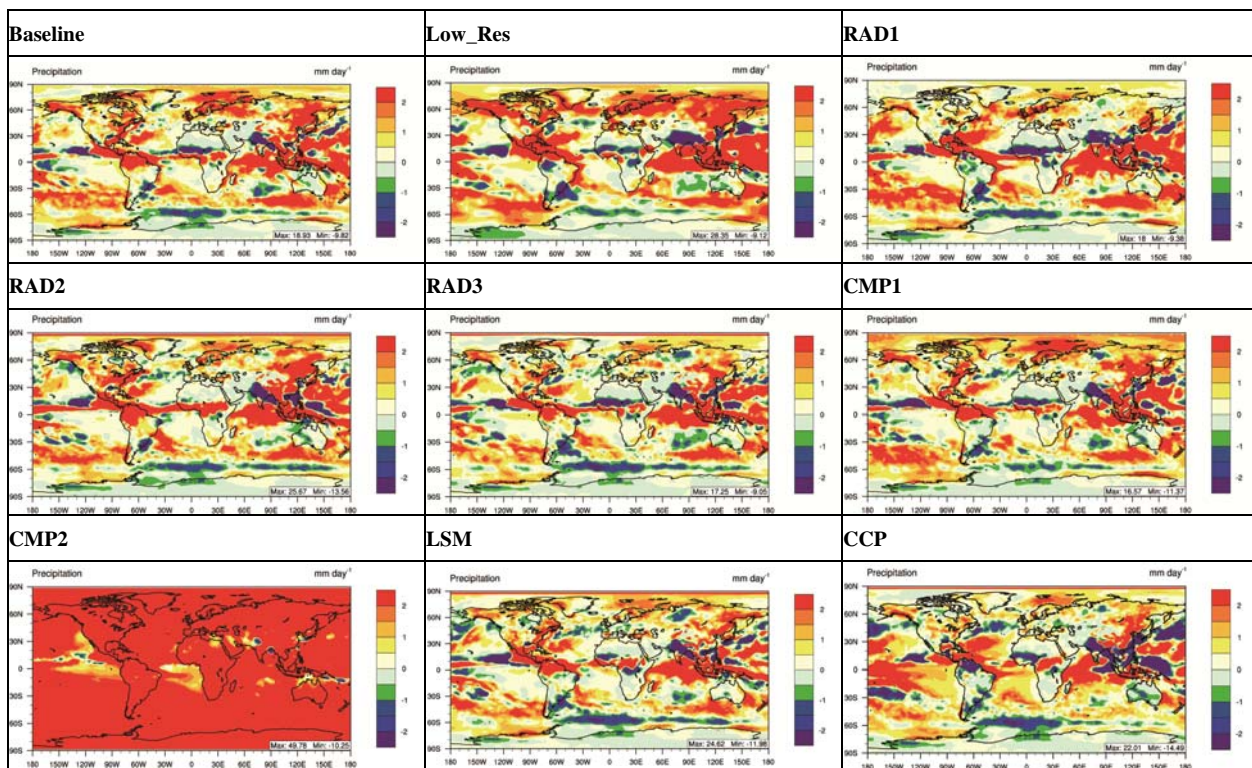


Figure 11. 2001 summer mean bias plot of daily mean precipitation rate against NNR data at a horizontal grid resolution of $1^\circ \times 1^\circ$ from the baseline, RAD1, RAD2, RAD4, CMP1, CMP2, LSM, and CCP simulations and at a horizontal grid resolution of $4^\circ \times 5^\circ$ from the Low_Res simulation.

Table 6. Performance statistics of GWRf baseline and sensitivity simulations for summer at $1^\circ \times 1^\circ$ on a global scale.

Variable	Dataset	Run	Mean Obs	Mean Mod	Number	Corr	MB	RMSE	NMB (%)	NME (%)
SW ($\text{W}\cdot\text{m}^{-2}$)	NNR	Base	191.34	198.94	61,560	0.93	7.59	41.34	3.97	15.08
		RAD1	191.34	159.61	61,560	0.88	-31.74	60.51	-16.59	23.42
		RAD2	191.34	200.68	61,560	0.95	9.34	34.74	4.88	12.53
		RAD3	191.34	194.71	61,560	0.95	3.37	35.66	1.76	12.85
		CMP1	191.34	186.97	61,560	0.91	-4.37	45.79	-2.28	17.26
		CMP2	191.34	161.60	61,560	0.92	-29.75	50.79	-15.55	19.60
		LSM	191.34	193.49	61,560	0.91	2.14	46.03	1.12	17.30
		CCP	191.34	164.26	61,560	0.91	-27.09	52.73	-14.16	20.09
	BSRN	Opt	191.34	152.94	61,560	0.91	-38.40	58.52	-20.07	22.65
		Base	218.83	270.32	29	0.98	51.50	57.07	23.53	23.53
		RAD1	218.83	226.99	29	0.98	8.16	21.08	3.73	8.19
		RAD2	218.83	268.67	29	0.97	49.85	54.95	22.78	22.78
		RAD3	218.83	261.92	29	0.98	43.09	48.51	19.69	19.69
		CMP1	218.83	259.31	29	0.98	40.48	46.18	18.50	18.70
LW ($\text{W}\cdot\text{m}^{-2}$)	NNR	CMP2	218.83	210.63	29	0.96	-8.20	24.55	-3.75	8.85
		LSM	218.83	272.21	29	0.97	53.38	60.70	24.40	24.40
		CCP	218.83	244.39	29	0.93	25.57	47.45	11.68	18.82
		Base	306.45	293.76	64,080	0.98	-12.68	24.30	-4.14	6.65
		RAD1	306.45	284.53	64,080	0.98	-21.92	29.67	-7.15	8.49
		RAD2	306.45	314.42	64,080	0.99	7.98	16.65	2.60	4.12
		RAD3	306.45	293.36	64,080	0.98	-13.08	23.61	-4.27	6.47
		CMP1	306.45	304.02	64,080	0.98	-2.42	23.38	-0.79	6.34
	BSRN	CMP2	306.45	336.00	64,080	0.97	29.55	51.81	9.64	11.66
		LSM	306.45	295.86	64,080	0.98	-10.59	23.88	-3.45	6.64
		CCP	306.45	306.73	64,080	0.98	0.28	21.18	0.09	5.80
		Opt	306.45	335.14	64,080	0.98	28.69	33.55	9.36	9.57
		Base	350.01	317.91	29	0.93	-32.10	39.63	-9.17	9.17
		RAD1	350.01	306.27	29	0.93	-43.74	49.36	-12.50	12.50
BSRN	RAD2	350.01	342.29	29	0.93	-7.72	25.11	-2.21	4.64	
	RAD3	350.01	317.48	29	0.92	-32.53	40.56	-9.29	9.33	
	CMP1	350.01	325.69	29	0.93	-24.32	32.81	-6.95	7.28	
	CMP2	350.01	358.23	29	0.92	8.22	28.07	2.35	5.74	
	LSM	350.01	316.50	29	0.92	-33.51	40.97	-9.57	9.73	
	CCP	350.01	323.15	29	0.92	-26.86	38.18	-7.67	8.76	

Continued

T2 (°C)	NNR	Base	7.48	7.68	64,079	0.99	0.20	2.33	2.70	18.75
		RAD1	7.48	6.31	64,079	0.99	-1.17	3.02	-15.64	27.34
		RAD2	7.48	8.08	64,079	0.99	0.60	2.47	8.01	20.61
		RAD3	7.48	7.49	64,079	0.99	0.01	2.28	0.14	18.84
		CMP1	7.48	7.45	64,079	0.99	-0.03	2.38	-0.38	20.58
		CMP2	7.48	11.77	64,078	0.96	4.29	10.59	57.30	62.08
		LSM	7.48	7.48	64,079	0.99	0.01	2.61	0.07	20.72
		CCP	7.48	6.74	64,079	0.99	-0.74	2.34	-9.92	20.65
	OPT	7.48	7.90	64,078	0.99	0.42	2.85	5.60	23.22	
	NCDC	Base	20.77	20.27	199	0.88	-0.51	4.18	-2.43	12.24
		RAD1	20.78	17.78	199	0.86	-3.00	5.33	-14.43	20.15
		RAD2	20.78	21.60	199	0.87	0.82	4.36	3.95	12.01
		RAD3	20.78	20.21	199	0.88	-0.57	4.18	-2.73	12.26
		CMP1	20.78	19.96	199	0.88	-0.81	4.21	-3.91	12.63
CMP2		20.78	20.74	199	0.86	-0.03	4.24	-0.17	11.49	
LSM		20.78	19.84	199	0.89	-0.94	4.04	-4.50	11.77	
CCP		20.78	19.41	199	0.88	-1.37	4.42	-6.58	14.05	
Q2 (g·kg ⁻¹)	NNR	Base	8.85	8.64	64,080	0.99	-0.21	1.03	-2.39	7.12
		RAD1	8.85	8.19	64,080	0.98	-0.66	1.46	-7.45	10.29
		RAD2	8.85	8.63	64,080	0.99	-0.22	0.85	-2.52	5.78
		RAD3	8.85	8.64	64,080	0.99	-0.21	1.04	-2.39	7.41
		CMP1	8.85	8.70	64,080	0.99	-0.15	1.00	-1.74	7.09
		CMP2	8.85	9.91	64,080	0.98	1.06	1.69	11.96	13.04
		LSM	8.85	9.17	64,080	0.98	0.31	1.27	3.52	7.86
		CCP	8.85	8.37	64,080	0.99	-0.49	1.07	-5.49	7.39
		Opt	8.85	9.16	64,080	0.97	0.30	1.46	3.42	8.29
Precip (mm·d ⁻¹)	GPCP	Base	2.22	3.08	10,363	0.83	0.86	1.96	38.64	54.4
		RAD1	2.22	3.05	10,363	0.79	0.83	2.13	37.15	55.11
		RAD2	2.22	3.22	10,362	0.74	0.99	2.59	44.62	63.86
		RAD3	2.22	2.87	10,361	0.79	0.65	1.94	29.12	50.79
		CMP1	2.22	3.00	10,363	0.80	0.77	1.91	34.77	53.57
		CMP2	2.22	9.66	10,363	0.60	7.44	8.18	334.31	335.37
		LSM	2.22	3.00	10,362	0.73	0.78	2.33	35.08	57.12
		CCP	2.22	2.72	10,363	0.73	0.49	1.78	22.10	50.87
	Opt	2.22	2.65	10,363	0.70	0.42	1.83	19.08	51.53	
	NCDC	Base	93.20	116.77	172	0.67	23.57	88.50	25.29	58.48
		RAD1	93.20	98.90	172	0.70	5.70	77.26	6.12	49.58
		RAD2	93.20	125.36	172	0.60	32.16	110.75	34.51	70.19
		RAD3	93.20	104.72	172	0.72	11.52	78.91	12.36	49.61
		CMP1	93.20	111.67	172	0.66	18.48	89.37	19.83	56.85
CMP2		91.20	284.69	172	0.60	191.50	219.92	205.48	213.91	
LSM		93.20	99.74	172	0.67	6.54	81.80	7.02	49.14	
CCP		93.20	93.86	172	0.54	0.66	96.05	0.71	54.95	

SW—Downward shortwave radiative flux at surface; LW—Downward longwave flux at surface; T2—Temperature at 2-m; Q2—Water vapor mixing ratio at 2-m; Precip—Daily mean precipitation rate.

within $\pm 5\%$ for BASE, RAD2, RAD3, and CMP1, indicating a similar performance by the simulations with the Goddard and CAM shortwave radiation schemes, or the RRTM and CAM longwave radiation schemes, or the WSM3 and WSM6 cloud microphysics schemes. Larger biases in SW by RAD1, CMP2, and CCP indicate a poor performance of the combination of the schemes that include the Dudhia shortwave scheme or the Purdue Lin cloud microphysics scheme, or the Grell-Devenyi cumulus parameterization, as compared with their respective schemes used in BASE. Compared with the BSRN data, RAD1 performs the best among all simulations and much better than BASE (with MB values of 8.2 vs. 51.5 $\text{W}\cdot\text{m}^{-2}$, and NMBs of 3.7% vs. 23.5%). CMP2 also performs well, with an NMB of -3.8% .

For LW, NMBs are within $\pm 10\%$ against the NNR data for all simulations. CCP, CMP1, and RAD2 show a better performance against the NNR data than other simulations, with NMBs of 0.1%, -0.8% , and 2.6%, respectively. Among all simulations, CMP2 gives the worst performance with an NMB of 9.6% and the largest biases occurring over Antarctic and Greenland as shown in **Figure 9**, indicating a high sensitivity of LW to the cloud microphysics over these areas. As shown in **Figure 9**, LW over most tropics and mid-latitude areas is under-predicted from all simulations except for RAD2 and CMP2. LW radiation flux over high latitudes is over-predicted except for Arctic from RAD1. RAD2 and CMP2 overpredict LW radiation flux over most of the domain. NMBs are also within $\pm 10\%$ against BSRN data for all simulations except for RAD1. RAD2 and CMP2 show better performance against BSRN than other simulations, with NMBs of 2.2% and 2.4%, respectively. BASE and RAD2 give MB values of -12.7 and $8.0 \text{ W}\cdot\text{m}^{-2}$ against NNR and -32.10 and $-7.72 \text{ W}\cdot\text{m}^{-2}$ against BSRN, respectively, indicating that the combination of schemes that include the RRTM longwave module is more accurate in simulating longwave radiation than that includes the CAM longwave module under summer conditions.

For T2, three simulations, LSM, RAD3, and CMP1 show better performance and CMP2 shows the worst performance against the NNR data. CMP2, BASE, and RAD3 show better performance and RAD1 shows the worst performance against the NCDC data. As shown in **Figures 9** and **10**, the large overpredictions of T2 by CMP2 and RAD2 are clearly related to overpredictions of LW from both simulations. Compared with results from BASE, LSM shows smaller overpredictions over land areas except for a few areas including the Greenland, Saudi Arabia, and Iran where LSM shows a larger under prediction than BASE. This indicates that the Slab LSM performs well under summer conditions, as compared with the NOAA LSM used in BASE. RAD3 and CMP1 show smaller overpredictions over several regions in

cluding North America and Russia, indicating a good performance by the CAM shortwave radiation scheme and by the WSM6 microphysics module. In contrast to all other simulations, RAD1 underpredicts T2 over most land areas (except for South Asia), particularly over Russia and northeastern China, which is likely caused by large underpredictions of SW radiation flux over high latitudes in Northern Hemisphere as shown in **Figure 8** and underpredictions of LW radiation flux over most of the domain except for South Pole and South Ferrel as shown in **Figure 9**. The worst performance in T2 predictions by CMP2 against the NNR data is caused by large overpredictions of LW radiation fluxes over the South Ferrel cell, South Pole, and Greenland. The errors over the poles are not reflected in the statistics of CMP2 against NCDC because there is no station in Greenland and only 3 out of the 199 stations are in Antarctica. These results indicate a poor performance in T2 predictions by the combination of the schemes that includes Dudhia shortwave radiation scheme (*i.e.* RAD1) or the Purdue Lin microphysics module (*i.e.* CMP2) under summer conditions.

Compared with BASE, For Q2, all NMBs are within $\pm 5\%$ against NNR data except for CMP2 and RAD1. For Precip, all simulations overpredict against the GPCP data, with the least overpredictions by CCP and RAD3 (NMBs of 22.1% and 29%, respectively) and the worst overpredictions by CMP2 (NMB of 334.3%). Compared with the NCDC data, Precip is also overpredicted, with a small overprediction by CCP, RAD1, and LSM (NMBs of 0.7%, 6.1%, and 7%, respectively) and the largest overprediction by CMP2 (NMB of 205.5%). As shown in **Figure 11**, overpredictions occur over nearly the entire domain by CMP2 and over most areas except for the South Ferrel cell and South Pole by other simulations. The poor performance of CMP2 indicates that the deficiency of the combination of schemes that include the Purdue-Lin microphysics parameterization in simulating precipitation under the summer conditions.

Given high sensitivity of the model predictions to various physical options, in particular, SW radiation scheme, cloud microphysics, and cumulus parameterization, an additional sensitivity simulation is performed using the best schemes or parameterizations identified through the above sensitivity simulations. The overall best schemes identified include the CAM shortwave scheme in terms of SW, T2, Q2, Precip (by comparing BASE, RAD1, and RAD3), the RRTM longwave scheme in terms of LW (by comparing RAD2 and RAD3), the WSM6 microphysics module in terms of SW, LW, T2, Q2, and Precip (by comparing BASE, CMP1, and CMP2), the Slab LSM (by comparing BASE and LSM), and the Grell-Devenyi cumulus parameterization in terms of Precip (comparing BASE and CCP). A sensitivity

simulation using these schemes is performed (referred to as OPT in **Table 6**). OPT shows the best performance in Precip among all simulations, with an NMB of 19.1 against GPCP (comparing to 22.1% - 334.3% by other simulations). However, given a non-linearity of the interactions among various schemes, OPT does not show the best performance for other meteorological variables. For example, it gives an NMB of -20.1% for SW, 9.4% for LW, 5.6% for T2, and 3.4% for Q2 against the NNR data.

6. Future Year Simulations

A simulation for the year of 2050 are performed at a horizontal grid resolution of $4^\circ \times 5^\circ$ and compared with the 2001 results from Low_Res. Under the SRES B1 climate scenario, global mean values of LW, T2, and Q2 increase and those of SW, wind speed at 10-m, PBL height, and Precip decrease. **Figure 12** shows the annual mean absolute changes in T2 and Precip between 2050 and 2001. In response to $\sim 51\%$ increases in CO_2 mixing ratios, T2 increases by up to 5.2°C over most areas in the global domain except for a few areas including the South Pole, South Ferrel cell, and Atlantic ocean where T2 may de-

crease by as much as 6.2°C , with a net increase of $\sim 0.2^\circ\text{C}$ on a global mean. The largest increases in T2 occur over southern South Africa, East Asia, southern North America, northern South America, and some areas in the North Ferrel cell. Precip in 2050 decreases over most areas, with the largest decrease of $8.2 \text{ mm}\cdot\text{day}^{-1}$ occurring over the tropics between $60^\circ\text{E} - 160^\circ\text{E}$. Precip increases over the southern China, central Africa, Australia, a large area over the Pacific ocean and some areas over the Atlantic ocean in the southern hemisphere. The net change is a decrease in global mean precipitation amount by $\sim 0.5 \text{ mm}\cdot\text{day}^{-1}$. The projected changes are generally consistent with other GCM predictions in terms of both magnitudes and spatial distributions (e.g. [62]).

7. Conclusions

The ability of GWRF in reproducing observations of major boundary layer meteorological variables is evaluated by comparing the 2001 model predictions with data from observational networks (e.g. NCDC and BSRN) and gridded Reanalysis data (e.g. NNR and GPCP). GWRF at a horizontal grid resolution of $1^\circ \times 1^\circ$ shows small biases against the NNR data and moderate over-

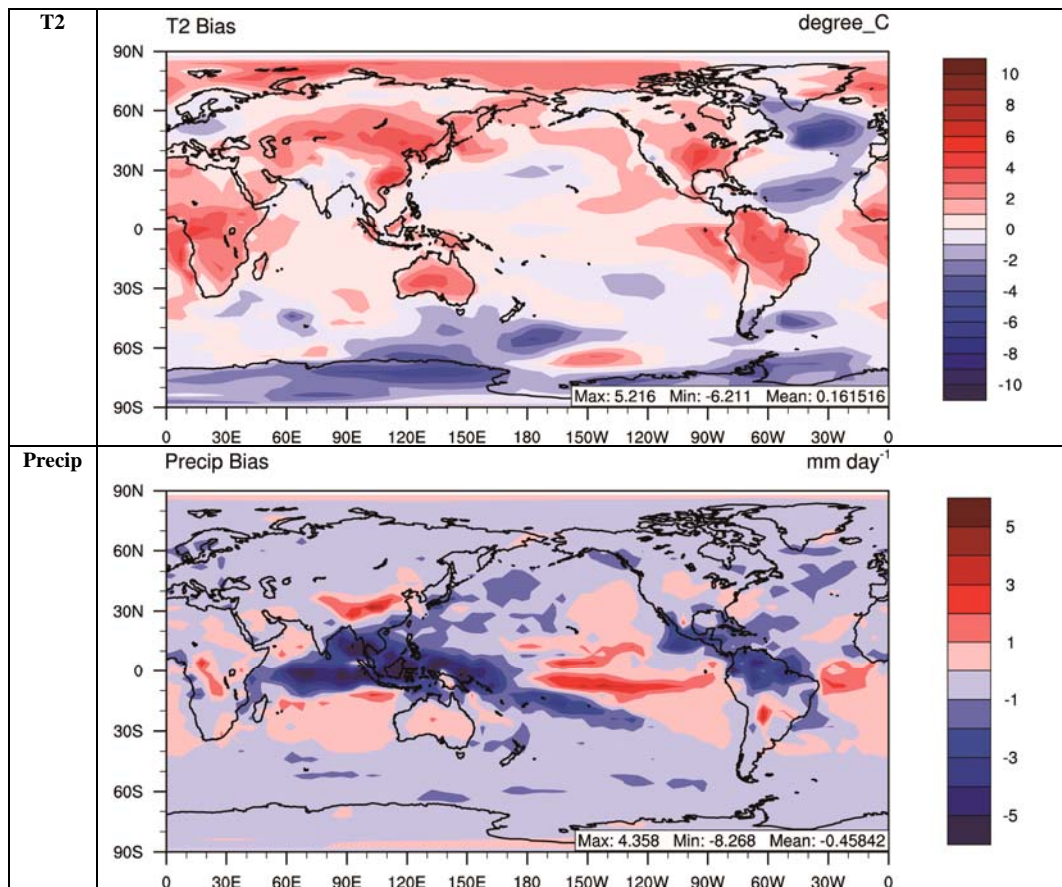


Figure 12. Projected absolute changes in temperature and precipitation in 2050 relative to 2001 at a horizontal grid resolution of $4^\circ \times 5^\circ$.

predictions against the BSRN data in the SW predictions. LW is slightly underpredicted against both the NNR and BSRN. T2 is underpredicted against the NNR and NCDC data in both summer and winter, with larger cold biases against the NCDC data (MBs as large as -2.3°C). Q2 is slightly underpredicted against the NNR data. Larger biases exist in wind speed and precipitation predictions, with large underpredictions in U10 against the NNR data and large overpredictions in Precip against the GPCP and NCDC data. The performance of GWRP in predicting major meteorological variables is overall consistent with that of most current GCMs.

Among all major variables examined, GWRP shows a large sensitivity to horizontal grid resolution for T2 predictions in winter and for Precip predictions in winter and summer. The model simulation at a coarser resolution leads to a much worse overprediction of T2 over the North America and a much less overprediction over the North Ferrel cell in winter. It also gives a worse agreement with observations for Precip over the most continents and most circulation cells during winter and summer. The sensitivity simulations using various physical options show a poor performance in SW, T2, and Q2 predictions by the combination of schemes that includes the Dudhia shortwave radiation scheme, or the Purdue Lin microphysics module, or the Grell-Devenyi cumulus parameterization, and in Precip predictions by the combination of schemes that includes the Purdue Lin microphysics module. These simulations identify that CAM, RRTM, WSM6, Slab, and Grell-Devenyi represent the best shortwave, longwave, cloud microphysics, land-surface, and cumulus treatments. While the simulation with these best schemes/parameterizations gives the best performance in Precip predictions by reducing the NMB of Precip from 38.6% to 19.1%, it does not give the best performance for all other major meteorological variables, because of a complex interaction of various atmospheric processes. A one-year simulation of GWRP for 2050 indicates a projected warmer and drier future climate. The projected future climate change may have important implications on the fate and lifetime of air pollutants. For example, decreased precipitation will reduce the scavenging rate of air pollutants, and decreased wind speeds and PBL height will increase the concentrations of air pollutant via reduced ventilation. The impact of future climate change on future air quality is being studied using a coupled GWRP and chemistry model and the results will be reported in a separate paper [63,64].

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